

## THE VERTICAL DISTRIBUTIONS OF ATMOSPHERIC $\text{CF}_2\text{Cl}_2$ , $\text{CFCl}_3$ AND $\text{N}_2\text{O}$ OVER SYOWA STATION IN 1983

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**Abstract:** Air samples were collected on an aircraft over Syowa Station up to an altitude of 7.3 km on January 24, April 5, October 18 and December 12, 1983. Trace gases ( $\text{CF}_2\text{Cl}_2$ ,  $\text{CFCl}_3$  and  $\text{N}_2\text{O}$ ) were measured by a GC-ECD method. These gases were vertically mixed well. Seasonal variation of vertical profiles could not be detected.

### 1. Introduction

There are only a few observations of vertical distributions of trace gases in Antarctica (GOLDAN *et al.*, 1980; ROBINSON *et al.*, 1982, 1983). GOLDAN *et al.* (1980) reported the vertical distributions of  $\text{CF}_2\text{Cl}_2$ ,  $\text{CFCl}_3$  and  $\text{N}_2\text{O}$  in the stratosphere, which were obtained by balloon-borne air samplings, and there was only one tropospheric air sample in each flight. Using an aircraft (C-130) and a helicopter, ROBINSON *et al.* (1982, 1983) observed higher mixing ratios of trace gases ( $\text{CF}_2\text{Cl}_2$ ,  $\text{CFCl}_3$ ,  $\text{N}_2\text{O}$ ,  $\text{CCl}_4$  and  $\text{CH}_3\text{CCl}_3$ ) in the upper troposphere than at the surface during the austral spring season. They suggested that these phenomena were brought about by the mid-latitude air mass advection over Antarctica through the upper troposphere.

In 1983, four flights were performed over Syowa Station, and air samples up to an altitude of 7.3 km were collected on "Pilatus PC-6" by the 24th Japanese Antarctic Research Expedition (JARE-24). In this paper, vertical distributions of tropospheric  $\text{CF}_2\text{Cl}_2$ ,  $\text{CFCl}_3$  and  $\text{N}_2\text{O}$  over Syowa Station will be reported.

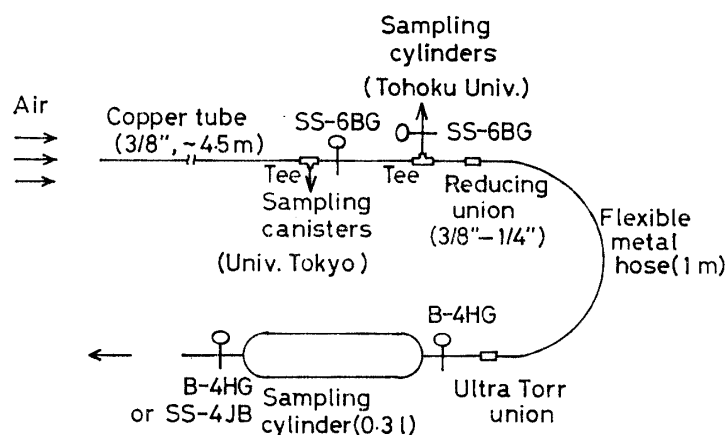


Fig. 1. Air sampling apparatus on the aircraft "Pilatus PC-6". After the first flight, sampling canisters (University of Tokyo) and SS-6BG valve between two Tees were removed.

## 2. Experimental

Air sampling method and gas-chromatographic measurement are described elsewhere (HIROTA *et al.*, 1984, 1985). On the aircraft, a copper tube (3/8") was installed in the stay of the wing in order to avoid the exhaust gas from the engine, and air was taken utilizing dynamic pressure as shown in Fig. 1 (HIROTA *et al.*, 1984).

After the first flight, sampling canisters (University of Tokyo) and SS-6BG valve between two Tees were removed.

As air samples were collected at atmospheric pressure, peak areas of gas chromatograms were corrected by the pressure of sample air.

## 3. Results and Discussion

Air samples were collected over Syowa Station (68.7–69.4°S, 39.4–40.2°E) up to the

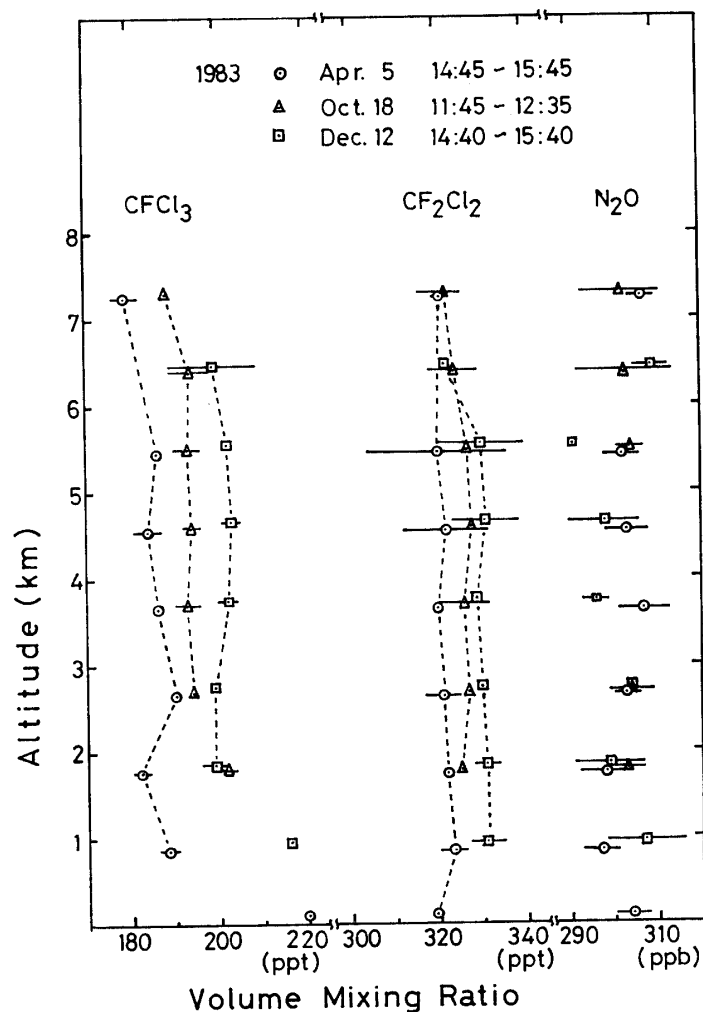


Fig. 2. Vertical distributions of CF<sub>2</sub>Cl<sub>2</sub>, CFCl<sub>3</sub> and N<sub>2</sub>O over Syowa Station. Differences of two measurements (>3 ppt for CF<sub>2</sub>Cl<sub>2</sub> and CFCl<sub>3</sub>, and >3 ppb for N<sub>2</sub>O) for each sample are shown by error bars.

Table 1. Mean volume mixing ratios of CF<sub>2</sub>Cl<sub>2</sub>, CFCI<sub>3</sub> and N<sub>2</sub>O over Syowa Station and their standard deviations\*.

Date	CF <sub>2</sub> Cl <sub>2</sub> (ppt)	CFCI <sub>3</sub> (ppt)	N <sub>2</sub> O(ppb)	Tropopause height** (km)
1983				
January 24	321(3)	180(3) <sup>a)</sup>	298(2)	8.6
April 5	321(1)	185(4) <sup>a)</sup>	303(4)	8.9
October 18	326(2)	193(2) <sup>b)</sup>	303(1)	12.2
December 12	329(3)	201(2) <sup>c)</sup>	301(6)	8.2

\* Values in parentheses. \*\* Sonde data at 1500 LT.

<sup>a)</sup> The value at 0.1 km was excluded.

<sup>b)</sup> The value at 1.8 km was excluded.

<sup>c)</sup> The value at 0.9 km was excluded.

altitude of 7.3 km on January 24, April 5, October 18 and December 12, 1983. Air samples of the first flight were carried back by JARE-23 prior to those of the later flights. The results of the first flight were reported in the previous paper (HIROTA *et al.*, 1984). Those of the second, third and fourth flights are shown in Fig. 2. Air sampling altitudes overlapped each other in all flights, data points are slightly shifting in Fig. 2. Differences of two measurements (> 3 ppt for CF<sub>2</sub>Cl<sub>2</sub> and CFCI<sub>3</sub>, and > 3 ppb for N<sub>2</sub>O) for each sample are indicated by error bars. Large error bars are generally seen in samples of higher altitudes. This would be due to smaller amounts of higher altitude samples according to lower atmospheric pressure. Mean values and standard deviations in each flight are shown in Table 1 together with the tropopause heights at 1500 LT.

High values of CFCI<sub>3</sub> in lower altitudes were excluded in the calculation of mean values (see footnotes of Table 1). These high values could not be expected from the contamination from Syowa Station, because high values were observed up to the altitude of 2 km and because high values of CF<sub>2</sub>Cl<sub>2</sub> were not observed from the same sample. As air samples were collected in the descent flight, it was impossible to take the contaminated air in the sampling tube into the sampling cylinder at lower altitude. These high values of CFCI<sub>3</sub> would be due to the adsorption of CFCI<sub>3</sub> on metal surface of the sampling tube, which was cooled below -35°C at about 6-7 km. With increasing temperature in the descent flight, CFCI<sub>3</sub> on metal surface would be desorbed and be taken into the sampling cylinder. In accordance with this speculation, these high values were observed only when the outer temperature was higher than -10°C. The lack of similar phenomena for CF<sub>2</sub>Cl<sub>2</sub> would be due to its lower boiling temperature than CFCI<sub>3</sub> by 54°C.

Mean values of CF<sub>2</sub>Cl<sub>2</sub> and CFCI<sub>3</sub> in Table 1 show that these gases were increasing during 1983 as was shown at the surface at Syowa Station (HIROTA *et al.*, 1985). That of N<sub>2</sub>O was almost constant within the range of experimental uncertainty.

Figure 2 shows that these gases were vertically well mixed up to the altitude of 6 km in spring, summer and autumn including the results on January 24. Above this level, CF<sub>2</sub>Cl<sub>2</sub> and CFCI<sub>3</sub> seemed to decrease slightly. This would be due to the lower tropopause heights on April 5 and December 12. On October 18, the tropopause height was 12.2 km. Its variation, however, was rather large and the mean value for

five days before and after the day was 10.5 km. Then a sudden increase of total ozone in this year was first observed on October 25 (ATMOSPHERIC ENVIRONMENT SERVICE OF CANADA, 1984). Generally the Antarctic tropopause is indistinct and frequently dissipates in winter and spring. Therefore, the tropospheric/stratospheric mixing might have been rather active around the middle of October 1983 before the sudden increase of total ozone.

ROBINSON *et al.* (1983) reported that higher mixing ratios of trace gases ( $\text{CF}_2\text{Cl}_2$ ,  $\text{CFCl}_3$ ,  $\text{N}_2\text{O}$ ,  $\text{CCl}_4$  and  $\text{CH}_3\text{CCl}_3$ ) in the upper troposphere and rather low values at the surface were observed during spring. They suggested that these phenomena were brought about by the mid-latitude air mass advection over Antarctica through the upper troposphere. Similar result was observed only for  $\text{CFCl}_3$  on December 12 over Syowa Station. This result, however, could not be significant considering the experimental uncertainty. In the observations of ROBINSON *et al.* (1983), statistically significant results were obtained only for  $\text{CF}_2\text{Cl}_2$  (Fig. 2 in ROBINSON *et al.*, 1983). In order to ascertain the long-range transports of trace gases over Antarctica, more precise measurement is needed.

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